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Albian high-resolution biostratigraphy and isotope stratigraphy: The Coppa della Nuvola pelagic succession of the Gargano Promontory (Southern Italy)

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Key words: Biostratigraphy, isotope stratigraphy, planktonic foraminifera, calcareous nannofossils, Albian, Marne a Fucoidi, Scaglia, Gargano Promontory

ABSTRACT

High-resolution $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ curves, calibrated against planktonic foraminiferal and calcareous nannofossil biostratigraphy, are provided for the upper Aptian–lower Cenomanian pelagic succession of the Gargano Promontory (Coppa della Nuvola section, southern Italy). The succession consists of two superimposed formations: the Marne a Fucoidi and the Scaglia (lower portion only). According to our integrated biostratigraphy, the entire succession spans the latest Aptian (planktonic foraminiferal *T. bejaouaensis* and calcareous nannofossil *R. angustus* Zones) and early Cenomanian stages (*R. cushmani* Zone CC9c). The Marne a Fucoidi–Scaglia transition falls in the late Albian (*R. ticinensis* Zone; CC9a+b Subzones).

The high-resolution $\delta^{13}\text{C}$ curve from the Coppa della Nuvola section can be subdivided into characteristic segments. Four negative shifts of $\delta^{13}\text{C}$ are recorded, followed by increasing values in, respectively, the early Albian (C11, C12), the early late Albian (C14, C15), the late Albian (C16, C17), and the early Cenomanian (C22). The late Albian carbon-isotope event, corresponding to the Oceanic Anoxic Event (OAE) 1d or Breistroffer Event, is possibly missing in the Coppa della Nuvola section as a result of condensation or erosion (C19–C20?). Even if, in the section studied, carbon-rich levels are not recorded (though some stratigraphic intervals are covered by Quaternary deposits in the lower portion), it is likely that the lower Albian (*T. primula*; *P. columnata* Zones) and the upper Albian (*T. praeticinensis* Subzone–*R. subticinensis* Zone; *R. achlyostaurion* Zone) positive $\delta^{13}\text{C}$ peaks succeeding negative trends in $\delta^{13}\text{C}$ (C11–C12; C16–17) record the pattern of global carbon burial, documented in other areas of the Gargano Promontory and elsewhere and connected to the OAE1b and OAE1c. Some or all of the negative $\delta^{13}\text{C}$ shifts may record the introduction of isotopically light carbon into the ocean–atmosphere system from the dissociation of gas hydrates.

The $\delta^{18}\text{O}$ curve of the Coppa della Nuvola section shows a similar trend to that of the $\delta^{13}\text{C}$ curve. Although partly of diagenetic origin, the negative shifts in the early Albian, late Albian and early Cenomanian may be interpreted as records of warming events resulting from the introduction of methane and its oxidation product carbon dioxide into the atmosphere.

The palaeoceanographic conditions inferred by the biotic and isotopic changes suggest fluctuating meso-eutrophic conditions through the late-Aptian–early-middle Albian and increasingly stable oligotrophic situation starting from the late Albian interval. Two main possible episodes of increased eutrophy, suggested by pulses in radiolarian abundance and marked drop in foraminiferal diversity, occurred during the Albian and correlate with the above-mentioned positive carbon-isotope shifts and regional and supraregional accumulation of organic matter (OAE1b and OAE1c).

RIASSUNTO

In questo lavoro sono presentate le curve ad alta risoluzione degli isotopi stabili del carbonio e dell'ossigeno ($\delta^{13}\text{C}$, $\delta^{18}\text{O}$), calibrate con la biostratigrafia a foraminiferi planctonici e nannofossili calcarei per l'intervallo cronostratigrafico Aptiano superiore–Cenomaniano inferiore della successione pelagica di Coppa della Nuvola (Promontorio del Gargano, Italia meridionale). Dal punto di vista litostratigrafico, la successione è costituita nella porzione inferiore dalle Marne a Fucoidi e nella porzione superiore dalla Scaglia. Lo studio biostratigrafico integrato della sezione ha permesso di riferire l'intera successione all'intervallo compreso tra l'Aptiano superiore [Zona a *Ticinella bejaouaensis* (foraminiferi planctonici) e Zona a *Rhagodiscus angustus* (nannofossili calcarei)] ed il Cenomaniano inferiore (Sottozona a nannofossili CC9c e Zona a foraminiferi *Rotalipora cushmani*). Il limite tra le Marne a Fucoidi e la Scaglia cade nell'Albiano superiore [Zona a *Rotalipora ticinensis* (foraminiferi planctonici); Sottozona CC9a+b (nannofossili calcarei)]. A differenza di altre successioni pelagiche del Gargano, nella sezione di Coppa della Nuvola non sono stati osservati livelli di *black shale*, anche se non è possibile escludere che questi coincidano con i tratti coperti nella parte inferiore della sezione.

La curva ad alta risoluzione del $\delta^{13}\text{C}$ può essere suddivisa in diversi segmenti caratteristici. In particolare, quattro picchi negativi, ognuno dei quali seguito da un aumento del valore, sono osservabili rispettivamente nell'Albiano inferiore (C11, C12), alla base dell'Albiano superiore (C14, C15), al termine dell'Albiano superiore (C16, C17) e nel Cenomaniano inferiore (C22). Tali picchi negativi sono separati da tre intervalli nei quali i valori del $\delta^{13}\text{C}$ si mantengono pressoché costanti. Sembra mancare, probabilmente per fenomeni d'erosione o condensazione, l'evento isotopico dell'Albiano superiore (C19–C20?) corrispondente all'evento anossico OAE1d (Evento Breistroffer). Nonostante nella sezione Coppa della Nuvola non siano stati rinvenuti livelli ricchi in carbonio organico, è comunque probabile che i picchi positivi del $\delta^{13}\text{C}$ documentati nell'Albiano inferiore (C12, Zona a *Ticinella primula* e a *Prediscosphaera columnata*) e nell'Albiano superiore (C16, Sottozona a *Ticinella praeticinensis* della Zona a *Rotalipora subticinensis* e Zona a *Rhagodiscus achlyostaurion*) corrispondano agli episodi globali di seppellimento di carbonio organico già segnalati per altre aree del Gargano e connessi con gli eventi anossici OAE1b e OAE1c. I picchi negativi nella curva del $\delta^{13}\text{C}$ possono rappresentare la documentazione dell'introduzione di grandi quantità di carbonio isotopicamente leggero nel sistema oceano-atmosfera in seguito a dissociazione di gas idrati.

La curva del $\delta^{18}\text{O}$ della sezione analizzata mostra un andamento simile alla curva del carbonio. Benché parzialmente di origine diagenetica, i picchi negativi osservati nell'Albiano inferiore, Albiano superiore e Cenomaniano inferiore possono essere interpretati come la registrazione di eventi di riscaldamento risultanti dall'introduzione di anidride carbonica come prodotto di ossidazione di metano.

Le condizioni paleoceanografiche, deducibili dai segnali biotici e isotopici, sono caratterizzate da fluttuazioni tra mesotrofia ed eutrofia durante l'intervallo compreso tra l'Aptiano superiore e l'Albiano inferiore e medio. A partire dall'Albiano superiore vengono ad instaurarsi condizioni più stabili tendenti all'oligotrofia. Due principali episodi di aumento dell'eutrofia, documentati durante l'Albiano, sono suggeriti da picchi d'abbondanza dei radiolari accompagnati da una marcata diminuzione della diversità dei foraminiferi. Tali episodi si correlano con i sopracitati picchi positivi degli isotopi del carbonio e accumulo regionale e supraregionale di sostanza organica (OAE1b e OAE1c).

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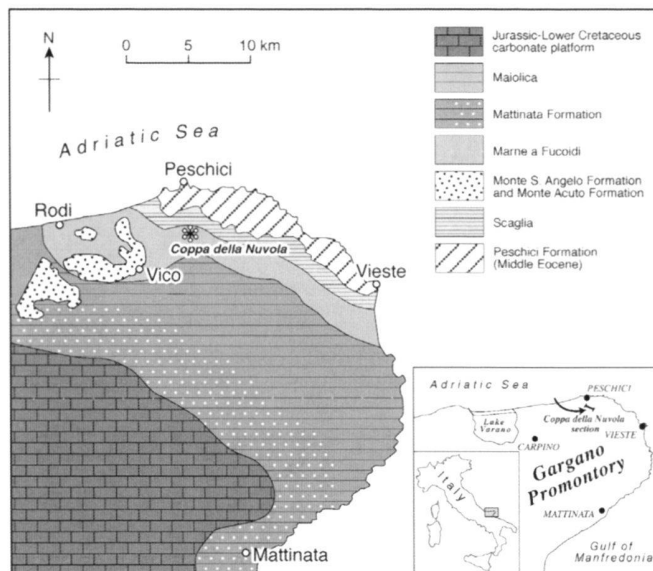


Fig. 1. Simplified geological and geographical map of the Gargano Promontory, showing the location of the section studied near the village of Peschici.

1. Introduction

The aim of this study is to describe biotic changes and accompanying carbon- and oxygen-isotope signals during the late Aptian–early Cenomanian, a time interval characterized by global palaeobiological and palaeoceanographic events. During this interval, a world-wide pulse or series of pulses in oceanic crustal production, together with formation of large igneous provinces, is credited with causing increased outgassing of carbon dioxide to produce a ‘greenhouse climate’ (Larson 1991; Tarduno et al. 1991; Jones & Jenkyns 2001). Warm temperatures and regional increases in primary productivity produced dysoxic and anoxic conditions on the sea floor and promoted several distinct episodes of organic-carbon accumulation (Oceanic Anoxic Events, OAEs: Schlanger & Jenkyns 1976; Jenkyns 1980; Arthur et al. 1990). The major oceanographic changes of this period also influenced microfossil evolution (e. g. Erbacher & Thürow 1997; Bischoff & Mutterlose 1998; Mutterlose 1998; Leckie et al. 2002).

To investigate the palaeoceanographic changes of the mid-Cretaceous, several hemipelagic and pelagic sections have been recently analysed using a combined biostratigraphical, sedimentological and geochemical approach. However, most of the studies to date have focused on the Aptian stage, which records the well-known Selli Event (OAE1a), whereas the Albian has attracted less attention. The pelagic succession cropping out in the northeastern part of the Gargano Promontory, southern Italy offers an almost continuous sedimentary record from the Hauterivian to the Coniacian stage, and this paper describes the upper Aptian–lower Cenomanian section at Coppa della Nuvola (Fig. 1).

The main goals of this study are: to provide high-resolution

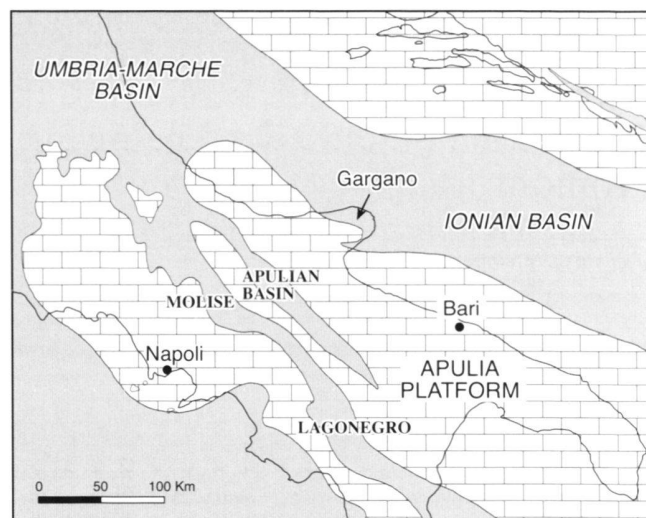


Fig. 2. Palaeogeography in southern Italy during the Jurassic–Cretaceous (after Zappaterra, 1990, modified)

$\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ curves, calibrated against planktonic foraminiferal and calcareous nannofossil data, in order to investigate respectively: 1) carbon cycling; 2) palaeotemperature evolution and (3) palaeoceanographic and biological changes through correlation between geochemical and micropalaeontological data.

2. Geological and stratigraphical setting

The Gargano Promontory is a peninsula almost entirely constituted by carbonate rocks belonging to the Apulian Platform in the southwest and the Ionian Basin in the northeast (Fig. 2). Modern research on the stratigraphy of the Gargano Promontory began in the mid-1960s by AGIP geologists and the Italian Geological Survey (Pavan & Pirini 1966; Martinis & Pavan 1967; Cremonini et al. 1971). These workers proposed a general stratigraphic framework, recognizing that the western half of the promontory was part of the shallow-water Apulia Platform while the eastern half was constituted by slope and basinal sediments. A series of biostratigraphic studies focused more recently on the slope and basinal deposits of the eastern Gargano Promontory (Luperto Sinni & Masse 1987; Coccioni & Luperto Sinni 1989; Luciani 1993; Luciani & Cobianchi 1994; Luperto Sinni & Borgomano 1994; Neri & Luciani 1994; Cobianchi et al. 1997; Cobianchi et al. 1999; Luciani et al. 2001). The current lithostratigraphic terminology of the Cretaceous basin-and-slope units varies according to different authors who have undertaken modern stratigraphical researches on the Gargano Promontory. The stratigraphical framework here adopted follows Bosellini et al. (1993), Cobianchi et al. (1997) and Bosellini et al. (1999).

The classical basinal Cretaceous succession recognized in the Umbria–Marche Basin, central Italy, can be extended far-

ther south to the Gargano Promontory, near the margin of the Apulia carbonate platform, though some lithological differences exist (Cobianchi et al. 1997; Cobianchi et al. 1999; Luciani et al. 2001). The Cretaceous pelagic succession across these areas consists of three superimposed stratigraphic units: Maiolica (Valanginian–early Aptian), Marne a Fucoidi (early Aptian–late Albian), and Scaglia (late Albian–Coniacian). Near the margin of the platform, the Mattinata Formation, rich in gravity-displaced deposits, laterally replaces the Maiolica (Bernoulli 1972; Bosellini et al. 1993; Luciani & Cobianchi 1994; Cobianchi et al. 1997; Graziano 2000).

The Apulian Platform, during the Late Jurassic–Early Cretaceous interval, was a productive carbonate platform with the rim colonized by bioconstructors and the inner area characterized by typical peritidal environments. The external margin zone passed gradually eastwards to the slope and finally to the basin (Morsilli & Bosellini 1997; Morsilli 1998). During the early Aptian, the Apulia Platform and the adjacent slope became abruptly inactive as a consequence of a drowning event (Bosellini et al. 1999; Cobianchi et al. 1997). During the latest Albian–Cenomanian interval, huge megabreccia bodies were deposited along the slope and base-of-slope belt coeval with the general emergence and karst development of the southern Apennine carbonate platforms.

Shallow-water deposits of late Aptian–Albian age are rare and crop out only locally (Bosellini et al. 1993; Bosellini et al. 2000). An inner-platform succession of Albian–Cenomanian age crops out in the southern Gargano and consists of mudstone–wackestone with peloidal packstone–grainstone intercalations containing orbitolinids of late Albian age (Merla et al. 1969; Luperto Sinni 1996). The late Aptian–Albian margin of the Apulia Platform is represented by bioconstructed limestones. Of this platform, only the upper Aptian portion is preserved (rudists, *Ellipsactinia* and stromatoporoids), whereas the Albian part is documented only in platform-derived breccias (rudists, corals, sponges and *Nerinea*) occurring in the slope and base-of-slope deposits (Luciani & Cobianchi 1994; Bosellini et al. 2000).

3. The Coppa della Nuvola section

3a. Location and lithology

The Coppa della Nuvola section, 106 m thick, is located southwest of the town of Peschici (Figs. 1, 3). In the lower part (from the base to 78 m) the succession consists of white fine-grained marly limestones, locally silicified (5–15 cm) interbedded with bioturbated grey marlstones (5–50 cm). The chert colour varies from whitish to brown or black. This unit can be attributed to the Marne a Fucoidi. Burrows, mainly represented by *Chondrites* and *Planolites*, also occur commonly in the marls and locally at the top of calcareous strata. The typical marl–limestone couplets of the Marne a Fucoidi formation are probably controlled by orbital (Milankovitch) parameters that influenced palaeoclimate and palaeoceanography (e. g., Premoli Silva et al. 1989a, b; Coccioni & Galeotti, 1991).

The upper portion of the section, referable to the Scaglia, is characterized by white, thinly bedded (10 cm) chalky lime mudstones with reddish chert, mainly in the form of nodules. The occurrence and thickness of marls decrease transitionally from the Marne a Fucoidi to the Scaglia, whereas the colour of the cherty nodules and layers changes sharply from the black-grey to reddish. Thin (10–15 cm) resedimented deposits, containing microbenthonic foraminifera, rare orbitolinid fragments, echinoids and bivalve fragments, occur locally in the Scaglia. Unlike other pelagic successions of the Gargano Promontory, no black shales have been observed in the Coppa della Nuvola section.

3b. Samples and data collection

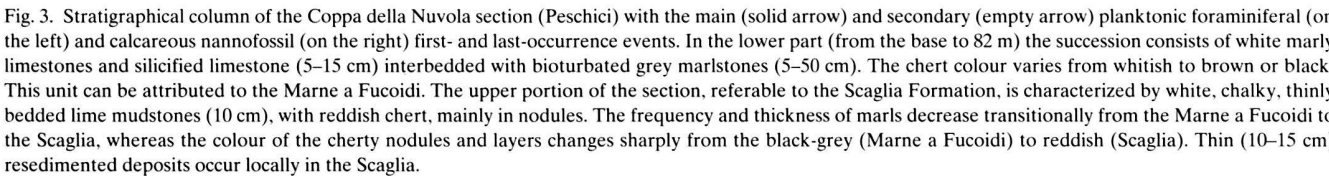
A total of 161 samples was taken at 50-cm intervals or less from the Coppa della Nuvola section. Samples were collected in order to carry out a detailed analysis of the calcareous nannofossil and foraminiferal content and to determine bulk carbon- and oxygen-isotope ratios. The same set of samples was used for micropalaeontological and geochemical studies in order to compare the results directly.

Standard laboratory methods for foraminiferal analysis were carried out: they consist of disaggregation of marly samples in Desogen and subsequent washing through a sieve of 38-micron screen in order to avoid the loss of very small specimens. The unevenly distributed planktonic foraminiferal assemblages and the different preservation throughout the section do not allow a quantitative analysis. 105 samples were taken from limestones and cherty limestones and analyzed in thin-section; the remaining 56 samples were analysed in washed residues.

Thin-sections were obtained also from the marlstones in order to evaluate throughout the section the radiolarian percentage abundances of the total sediment, according to the diagrams for the visual percentage estimation in sedimentary rocks given by Baccelle & Bosellini (1965).

For the calcareous-nannofossil analysis, smear-slides from samples of all the rock-types (marly limestone, marl, calcareous shale) were prepared and were analysed under a polarizing light microscope at a magnification of $\times 1250$. For each smear slide, 300 fields of view were observed in random traverses. Calcareous nannofossil totals and species were semi-quantitatively estimated.

Bulk samples for isotopic analysis were first powdered, cleaned with 10% H_2O_2 followed by acetone, and then dried at 60°C. Powders were then reacted with purified orthophosphoric acid at 90°C and analysed on-line using a VG Isocarb device and Prism Mass Spectrometer at Oxford University. Normal corrections were applied and the results are reported, using the usual δ notation, in per mil deviation from the PDB standard. Calibration to PDB was performed via our laboratory Carrara marble standard. Reproducibility of replicate analyses of standards was generally better than 0.1‰ for both carbon- and oxygen-isotope ratios.



In the following sections the main microfossil events are summarized.

Preservation and abundance of planktonic foraminiferal assemblages from the 161 samples analysed are highly variable, both in washed residue and in thin section, with two main intervals of poor preservation (from 16 to 26 m and from 66 to 70 m from the base). In spite of this, several planktonic foraminiferal events were recognised in the Coppa della Nuvola section. In particular the main events (Fig. 3) from the base are: the first occurrences of *T. primula* (sample 20, 11 m), *Biticinella breggiensis* (sample 48, 29.2 m), *Rotalipora subticinensis* (sample 88, 71 m), *Rotalipora ticinensis* (sample 98, 76m), *Rotalipora appenninica* (sample 117, 85.70m), *Rotalipora globotruncanoides* (sample 124, 89.5m), *Rotalipora reicheli* (sample 153, 104.6m), *Rotalipora cushmani* (sample 155, 105.5m). Low-latitude planktonic foraminiferal zonation for the mid-Cretaceous is well established and reliable and the different zonal schemes proposed are rather similar (e. g. Caron 1985; Sliter 1989a, 1992; Premoli Silva & Sliter 1994; Robaszynski & Caron 1995). On the basis of the events observed in the studied section, the following biozones were identified: *Ticinella bejaouaensis* Zone, *Ticinella primula* Zone, *Biticinella breggiensis* Zone, (*Ticinella praeticinensis* Subzone, *Rotalipora subticinensis* Subzone), *Rotalipora ticinensis* Zone, *Rotalipora appenninica* Zone, *Rotalipora globotruncanoides* Zone, *Rotalipora reicheli* Zone, *Rotalipora cushmani* Zone. Beside the main events, the high-resolution analysis has enabled us to identify several secondary biostratigraphic datum levels, which are shown in Fig. 3.

For this study 156 samples were analysed for their constituent calcareous nannofossils. Nannofloral abundance ranges from “few” to “abundant” with minor modification due to the effects of dissolution and calcite overgrowth. In the studied section several nannofossil events were recognized, some of which have been used to define zonal or subzonal boundaries by previous authors (Thierstein 1971, 1973, 1976; Manivit et al. 1977; Sissingh 1977; Perch-Nielsen 1985; Bralower 1987; Erba 1988; Ghisletti & Erba in Premoli Silva & Sliter 1994; Bralower et al. 1995; Erba 1996; Bown et al. 1998; Erba et al. 1999; Premoli Silva et al. 1999; Bralower et al. 1999). These events are plotted in figure 3 and listed below, along with some additional events (from bottom to top): the LO (last occurrence) of *Assipetra infracretacea* at the bottom of the measured section; the FO (first occurrence) of *Prediscosphaera columnata* at 6 m from the bottom of the section; the FO of *Cribrosphaerella ehrenbergii* at 7 m; the FO of *Axopodorhabdus albianus* at 25 m; the FO of *Biscutum magnum* at 28.5 m; the FO of *Rhagodiscus achlyostaurion* at 32m; the FO of the genus *Eiffel-*

Based on these events, five biozones have been identified, from older to younger: the *Rhagodiscus angustus* Zone (Thierstein 1973; 0–6 m); the *Prediscosphaera columnata* Zone (Hill 1976; 6–25 m); the *Axopodorhabdus albianus* Zone (Hill, 1976, emended by Erba 1988; 25–32 m); the *Rhagodiscus achlyostaurion* Zone (Erba 1988; 32–75.5 m) and the *Eiffelithus turriseiffelii* Zone (Thierstein 1971, emended by Sissingh 1977; 73 m–top of the section).

The nannofossil assemblages are diversified and moderately well preserved. Species abundant in this interval include: *Watznaueria barnesae*, *Cretarhabdus surirellus*, *Lithraphidites carniolensis*, *Manivitella pemmatoidea*, *Microstaurus chistiatus*, *Rhagodiscus asper*, *Rhagodiscus splendens*, *Rucinolithus irregularis*, *Vagalapilla stradneri*, *Zeughrabdothus embergeri*. Nanoconids are very rare, discontinuously recorded and mainly represented by *Nannoconus truittii* and *N. elongatus*.

Figure 4 shows the main biostratigraphic events recorded in this study and correlates the calcareous nannofossil and planktonic foraminiferal biozones with mid-Cretaceous stages. The integrated biostratigraphic scheme proposed here largely confirms the standard Tethyan integrated biostratigraphy (see Hardenbol et al. 1998); the main differences with previous schemes are discussed below.

1 – The foraminiferal *Hedbergella planispira* Zone, generally placed between the *Ticinella bejaouaensis* and *Ticinella primula* Zones (e. g. Sliter 1989a; Robaszynski & Caron 1995), was not identified in the Coppa della Nuvola section, a circumstance already recorded for other sections of the Gargano Promontory (Cobianchi et al. 1997). Therefore, the foraminiferal *Ticinella bejaouaensis* Zone correlates with both the nanofossil *Rhagodiscus angustus* Zone and the lower portion of the *Prediscosphaera columnata* Zone.

According to the definition of Caron (1985), the *Hedbergella planispira* Zone represents the interval from the appearance of the zonal marker to the first occurrence of *Ticinella primula*. In many localities the upper boundary corresponds, however, to an interval characterized by a low-diversity assemblage dominated by *Hedbergella planispira* including several levels that are completely barren. Such assemblages are characteristically found in black shales (e. g. Br  h  ret et al. 1986; Premoli Silva et al. 1989 a, b; Sliter 1989a, b). The absence of larger taxa (ticinellids) has been interpreted as a response to an expanded oxygen-minimum zone that destroyed the habitats of the relatively larger, less tolerant forms. The barren levels can also be the result of intense dissolution at the sediment-water interface. The depauperate interval was attributed to the *Hedbergella planispira* Zone; both the upper and lower boundaries of the *Hedbergella planispira* Zone are

Fm	CALCAREOUS NANNOFOSSILS		PLANKTONIC FORAMINIFERA		AGE
SCAGLIA	<i>C. kennedyi</i>	c	<i>R. cushmani</i>	<i>R. cushmani</i>	CENOM.
			<i>R. reicheli</i>	<i>R. reicheli</i>	
MARNE A FUCOIDI	<i>E. turriseiffelii</i>	a + b	<i>R. globotruncanoides</i>	<i>R. globotrunc.</i>	—98.9—
			<i>R. appenninica</i>	<i>R. appenninica</i>	
			<i>R. ticinensis</i>	<i>R. ticinensis</i>	late
			<i>R. subtic.</i>	<i>R. subticinensis</i>	
		<i>R. achlyostaurion</i>	<i>R. praetic.</i>		middle
		<i>A. albianus</i>	<i>B. breggiensis</i>	<i>B. breggiensis</i>	
		<i>A. albianus</i>			early
		<i>P. columnata</i>	<i>T. primula</i>	<i>T. primula</i>	
		<i>P. columnata</i>			—112.2—
		<i>R. angustus</i>	<i>T. bejaouaensis</i>		
					late
					APTIAN

Fig. 4. Planktonic foraminiferal and calcareous nannofossil integrated biostratigraphy from the Coppa della Nuvola section (Peschici) with the events characterizing the zonal boundaries and correlation with formations and ages. The integrated biostratigraphical scheme mainly confirms the standard Tethyan integrated biostratigraphy, except for the absence of the *Hedbergella planispira* Zone. This zone, generally placed between the *T. bejaouaensis* and *T. primula* Zones, was not found in the Coppa della Nuvola section, as is also the case for the Ischitella section, cropping out south of Rodi Garganico (Cobianchi et al. 1997).

hence uncertain, as is the age and duration of the zone. It appears probable that the lack of black shale (implying probable absence of hostile conditions) in this interval of the Gargano succession, may explain the absence of the *Hedbergella planispira* zone, even though condensation or erosion cannot be excluded.

2 – In the scheme of Hardenbol et al. (1998), the FO of *Lithraphidites acutum*, which defines the boundary between the nannofossil CC9 and CC10 Zones (Manivit et al. 1977; Sissingh 1977), is placed in the foraminiferal *Rotalipora reicheli* Zone. In the Coppa della Nuvola section the FO of *L. acutum* was not detected, so the CC9c Subzone is taken to correlate with the upper part of the *Rotalipora globotruncanoides* Zone, the *R. reicheli* Zone and the lower portion of the *R.*

cushmani Zone. However, the species *L. acutum* is rare and discontinuously distributed in other Tethyan sections (Ghisletti & Erba in Premoli Silva & Sliter 1994; Luciani & Cobianchi 1999). Moreover, partial dolomitization occurring in the uppermost portion of the studied section has probably influenced the nannofossil preservation.

Owing to the absence of ammonites in the Coppa della Nuvola section, the correlation between biostratigraphy and chronostratigraphy of the mid-Cretaceous is indirect and based on schemes proposed by various authors (e.g. Erba 1988; Ghisletti & Erba in Premoli Silva & Sliter 1994; Bralower et al. 1995; Robaszynski & Caron 1995; Bellanca et al. 1996; Gale et al. 1996; Erba et al. 1999; Kennedy et al. 2000). The Marne a Fucoidi/Scaglia transition lies within the nannofossil *Eiffel-*

lithus turrisiiffeli Zone and within the upper part of the foraminiferal *Rotalipora ticinensis* Zone of late Albian age, as also reported for the Bottaccione section from the Umbria–Marche Basin (Premoli Silva & Sliter 1994). Therefore, the Marne or Scisti a Fucoidi (lower 78 m) of the Coppa della Nuvola section correspond to the late Aptian–late Albian, whereas the Scaglia (upper 26.25 m) is late Albian–early Cenomanian in age.

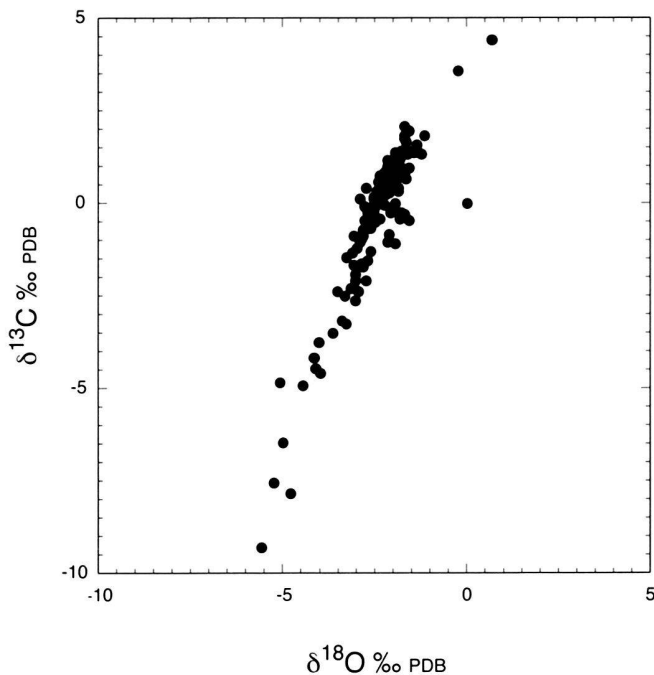


Fig. 5. Cross-plot of $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of bulk pelagic carbonates from the Coppa della Nuvola section. Note the wide spread of values and the apparent 'mixing line' produced by variable addition of a diagenetic cement to a primary marine carbonate. Such relationships may also have a primary origin, being produced by injection of dissociated gas hydrates ($\delta^{13}\text{C} \approx -60\text{‰}$) into the ocean-atmosphere system and consequent rapid global warming causing a fall in $\delta^{18}\text{O}$ values in skeletal and inorganic carbonate, followed by subsequent decay to equilibrium values of both isotopic parameters.

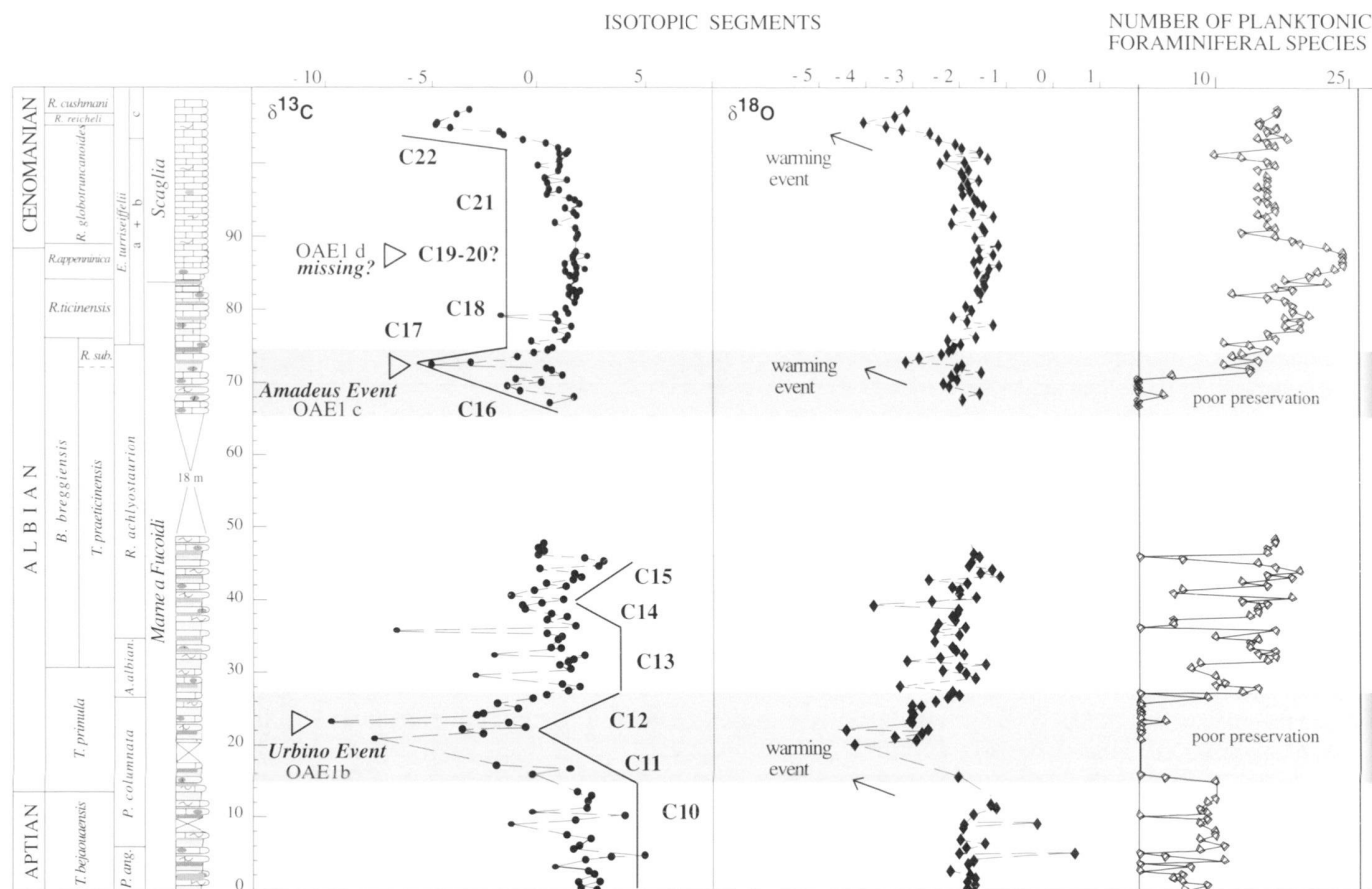


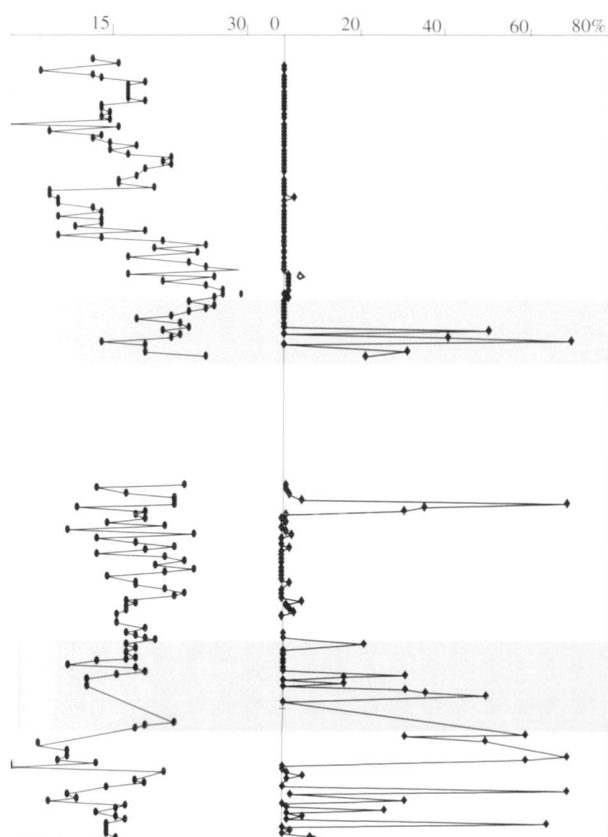
Fig. 6. Scheme showing the correlation between lithostratigraphy, biostratigraphy (foraminifera and calcareous nannofossils), stable carbon- and oxygen-isotope curves and planktonic foraminiferal and calcareous nannofossil species diversity and radiolarian abundance from the Coppa della Nuvola section. Note that the two isotopic curves display similar trends. We have produced a tentative correlation between the C10–C15 $\delta^{13}\text{C}$ isotopic segments identified by Bralower et al. (1999) from the Albian of the Santa Rosa Canyon section in Mexico and those from Peschici. The negative shift at the top of the section (C22; *R. reicheli*–*E. turrisaiffellii* Zones) probably correlates with a similar but more poorly defined feature recorded in the Contessa section (Gubbio, Umbria–Marche Basin; Stoll & Schrag 2000). A reliable correlation exists between the Gargano Promontory, the Vocontian Basin (Southeastern France) and the Magazan Plateau (Northwest Africa, DSDP Site 545) (Herrle, 2002) in the record of a marked $\delta^{13}\text{C}$ negative trend, followed by a positive shift, in the early Albian *Prediscosphaera columnata* Zone. This chemostratigraphic event occurred just before or during during a time of widespread anoxia and regional deposition of black shales (OAE1b). Possible warming and cooling events are underlined in the oxygen-isotope curve. A possible correlation between Albian carbon-isotope fluctuations, regional deposition of organic matter (Amadeus Segment, Urbino Level) and OAEs is also indicated.

their counterparts in Marche–Umbria, which contain more clay and are vividly coloured in shades of red, green and black: such differences are compatible with the hypothesis that considerable import of off-platform calcareous fines characterized deposition in the basin. Such reactive material may well have produced early to late diagenetic cements, whose isotopic composition should nonetheless have broadly reflected global trends with respect to carbon-isotope signatures.

As far as absolute values of the carbon-isotope profile are concerned, $\delta^{13}\text{C}$ values are both more scattered and lower by some 1–2‰ than those recorded from many Tethyan sections, including those from Marche–Umbria (Scholle & Arthur 1980; Bralower et al. 1999; Jenkyns et al. 1994; Stoll & Schrag 2000; Strasser et al. 2001). Such relatively low values were also

recorded for the Aptian interval of the Gargano Promontory (Luciani et al. 2001). In this case, the lower values were related to the possible input of fine-grained aragonitic and calcitic detritus of shallow-water origin derived from the Apulia Platform, where more restricted lagoonal environments were characterized by waters whose carbon-isotopic composition was influenced by oxidation of organic matter (cf. Patterson & Walter 1994). In fact, carbon-isotope values in low-latitude Cretaceous Tethyan platform carbonates are typically, but not invariably offset negatively by about -1‰ with respect to coeval pelagic facies, even though they follow global trends (Grötsch et al. 1998; Davey & Jenkyns 1999; cf. Ferreri et al. 1997). Additional support for the presence of a primary component in the isotopic signatures of the Coppa della Nuvola

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succession comes from the fact that the negative carbon-isotope excursion, stratigraphically registered close to the first occurrence of *A. albianus* (OAE 1b), is also recorded from northeastern Mexico in material other than carbonate, namely marine organic matter (Bralower et al. 1999). Given that the negative carbon-isotope excursion is additionally recognized in bulk sediments from the eastern Atlantic (Mazagan Plateau) and southern France, in close stratigraphic proximity to the Paquier black shale, the $\delta^{13}\text{C}$ signal is clearly a regionally reproducible phenomenon (Herrle 2002; Jenkyns 2003).

Primary palaeotemperature signals must be encoded in the oxygen-isotope record but cementation and/or recrystallization close to or below the sea floor may have grossly compromised the primary signatures of the planktonic component. However, given that the $\delta^{18}\text{O}$ data from the Atlantic site show up to a -1‰ negative shift in concert with the abrupt change to lower carbon-isotope values, and thus conform roughly to the isotopic profiles from Coppa della Nuvola, the possible primary nature of some of the signal also needs to be explored. For example, introduction of isotopically light carbon into the ocean–

atmosphere system, coincident with global temperature rise, would produce highly correlated negative $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ excursions that are totally unrelated to diagenesis.

5b. Isotopic curves

There are few high-resolution isotopic curves for the Albian and they generally describe only a portion of the stage. The detailed chemostratigraphical analysis of the Coppa della Nuvola section allowed us to reconstruct carbonate carbon- and oxygen-isotope curves for the Upper Aptian–Lower Cenomanian interval, and this represents one of the most detailed and complete isotopic carbonate records from low latitudes.

The $\delta^{13}\text{C}$ curve can be split into characteristic segments (Fig. 6) with partial reference to those documented by Bralower et al. (1999) from northeastern Mexico. The upper Aptian–lower Albian is characterized by relatively constant values of $\delta^{13}\text{C}$ (~ 1.8‰; segment C10). The $\delta^{13}\text{C}$ values then show an abrupt negative shift (C11) in the upper lower Albian (down to -3‰), followed by a return, in the middle Albian, to values of +1.8‰ (C12). Subsequently, the $\delta^{13}\text{C}$ values remain within the range of +0.5/+2‰ (C13), with a minor negative (-1‰; C14) and subsequent positive (+2‰; C15) excursion occurring in the lower upper Albian. A covered tract in the section, interrupting the isotopic record for part of the upper Albian interval, precedes a drop where values fall to -2.6‰ (C16). The following C17 segment represents a return to more positive values leading to a relatively constant tract (~ +0.8‰) spanning the uppermost Albian and the lower Cenomanian interval. Finally, the $\delta^{13}\text{C}$ curve defines a further negative spike recording the lowest values measured in this section (C21–22).

The Coppa della Nuvola $\delta^{13}\text{C}$ curve shows similar features to the $\delta^{13}\text{C}$ curve provided for Mexico (Scholle & Arthur 1980; Bralower et al. 1999), where equivalent C10–C15 segments have been recognized. Minor discrepancies in the biostratigraphical calibration can be explained by low biostratigraphical resolution for the Mexican sections due to poor general preservation of both calcareous nannofossil and planktonic-foraminiferal assemblages. Other analogies can be found with the $\delta^{13}\text{C}$ curves illustrated by Jenkyns et al. (1994) and Stoll & Schrag (2000) for sections through the Bottaccione and Contessa Gorges (Umbria–Marche Basin, Central Italy), where a negative shift (albeit more gradual with respect to the C22 Segment of the Coppa della Nuvola section) is recorded from the *Rotalipora globotruncanoides* Zone culminating in the *Rotalipora reicheli* Zone.

Pronounced positive shifts in Cretaceous carbon-isotope curves generally correlate with oceanic anoxic events (e. g. Jenkyns 1980, 1999; Scholle & Arthur 1980; Schlanger et al. 1987; Arthur et al. 1988; Menegatti et al. 1998; Larson & Erba 1999; Stoll & Schrag 2000), which record dramatic perturbations in the global carbon cycle, characterized by increases in marine productivity and burial rates of organic carbon. These processes are particularly evident in the case of OAE 1a (early Aptian Selli Event) and OAE 2 (Cenomanian–Turonian Bonarelli

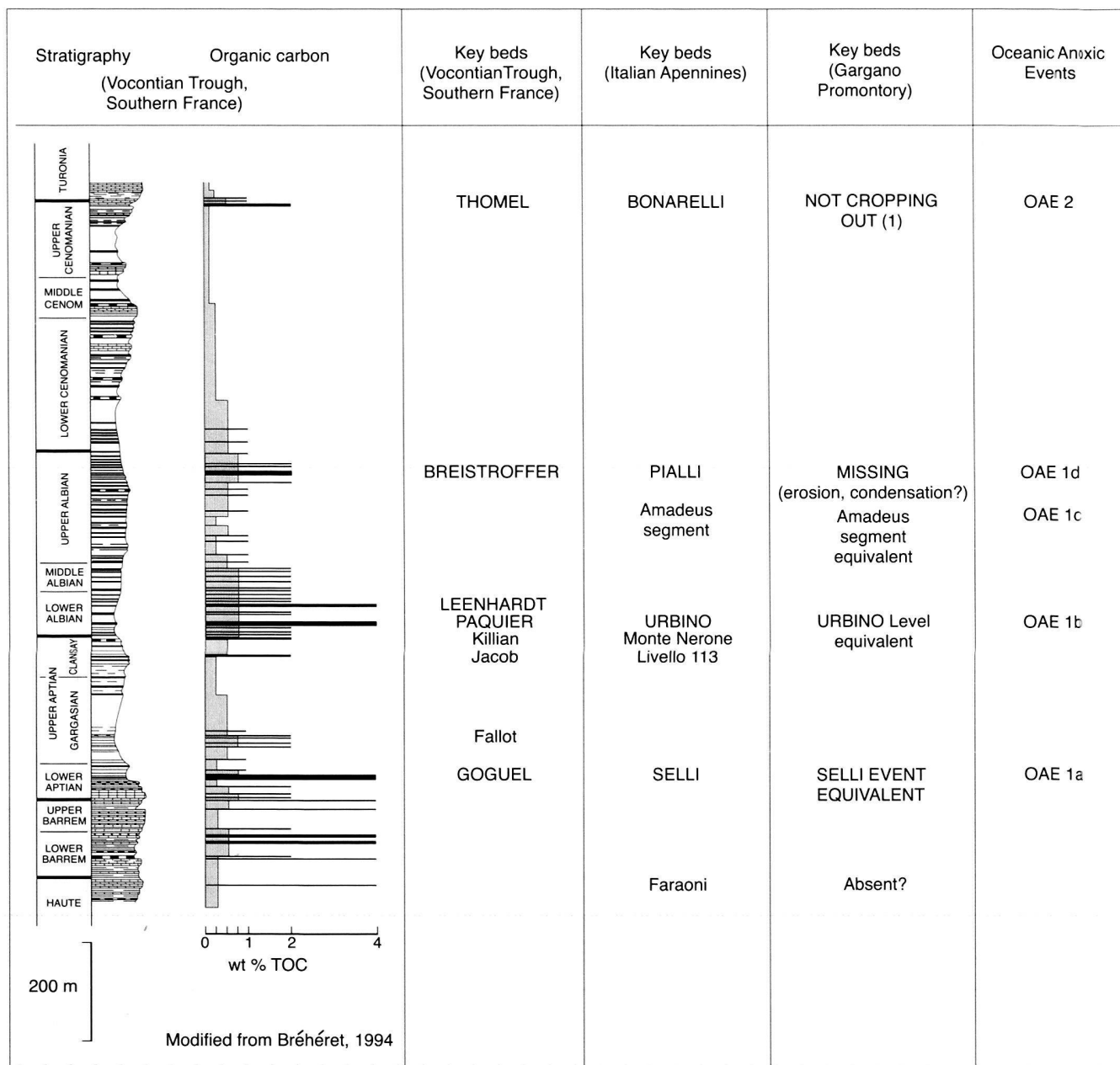


Fig. 7. Major organic-rich levels in the Vocontian Trough (southeastern France) and corresponding Oceanic Anoxic Events plotted against key beds from the Italian Apennines and Gargano Promontory. Note that the Bonarelli Level is recorded offshore from the Gargano Promontory (AGIP wells). Modified from Bréhéret (1994), with additional data from Arthur et al. (1990), Erbacher (1994), Coccioni (2001) and our observations.

Event). Complications arise in the case of the former, because of the presence of a pronounced negative carbon-isotope excursion registered immediately below the Selli Level and its equivalents (Jenkyns 1995; Menegatti et al. 1998; Jenkyns & Wilson 1999; Bralower et al. 1999; Erba et al. 1999; Erbacher et al. 1999; Luciani et al. 2001; Bellanca et al. 2002). Marked rapid negative excursions in carbon-isotope curves have been recently attributed to release of isotopically very negative ($\delta^{13}\text{C} \approx -60\text{‰}$)

methane deriving from dissociation of gas hydrates within continental-margin sediments (e.g. Dickens et al. 1995; Hesselbo et al. 2000; Jahren et al. 2001; Padden et al. 2001). The release of methane gas hydrate is an abrupt event and, because the gas undergoes near-instantaneous oxidation to CO_2 , it is efficiently dispersed into the ocean-atmosphere system. Rapid (10^3 yrs: Röhl et al. 2000) global warming ensues, registered by lowering of $\delta^{18}\text{O}$ values in biogenic carbonate (Jenkyns 2003).

In the Umbria–Marche Basin, as well as in the Swiss Prealps (Rotter Sattel section), no major carbon-isotope shifts have been detected around the beds recording OAE1b (Stoll & Schrag 2000; Strasser et al. 2001). However, the regional negative carbon-isotope shift recorded elsewhere suggests a global phenomenon that could be attributed to release of gas hydrates (Jenkyns 2003). This record, at or just below the level of the black shale (Paquier Level and equivalents: Fig. 7), is known from carbonate materials in the Vocontian Trough (Southeast France) and the Mazagan Plateau, Atlantic Ocean (Herrle 2002), and from northeast Mexico in organic matter (Bralower et al. 1999). The $\delta^{13}\text{C}_{\text{org}}$ record of bulk organic matter and certain biomarkers of the black shale itself, however, show a substantial positive shift ($\sim 8\text{--}10\text{‰}$), attributed to a fundamental change in the source of the organic matter over the interval in question (Kuypers et al. 2001). The exact nature of the negative $\delta^{13}\text{C}$ excursions associated with other Aptian–Albian anoxic events (OAE1c; Segment C16–17) as well as Segment C22 in the Cenomanian remains problematic, but dissociation of gas hydrates may also be involved in these instances.

The black shales recorded within the upper Albian *R. appenninica* Zone in several areas (Fig. 7), named as ‘Breistroffer Level’ and ‘Pialli Level’ (e. g. Northern Atlantic and western Tethys, Erbacher & Thurow 1997; Vocontian Basin, Erbacher et al. 1999; Swiss Prealps, Strasser et al. 2001; Umbria–Marche Basin, Coccioni 2001) and considered as the expression of the OAE1d, are not present in the Coppa della Nuvola section. The carbon-isotope curve does not display peaks across this interval; on the contrary there is a tract with relatively constant values (Fig. 6), as is also the case in the Roter Sattel section from the Swiss Prealps (Strasser et al. 2001). The $\delta^{13}\text{C}$ record is different in the tropical western Atlantic Ocean where a pronounced negative shift followed by a positive fluctuation marks the black-shale deposition characteristic of OAE1d (Wilson & Norris 2001). Similar isotopic signatures are also registered in the Vocontian Trough in southeastern France (Gale et al. 1996) and in the UK (Speeton, Yorkshire: Mitchell et al. 1996) but cannot be positively recognized in Marche–Umbria (cf. Stoll & Schrag 2000). The negative excursion in the early Cenomanian (Segment C22; Fig. 6) may be tentatively recognized in the $\delta^{13}\text{C}$ profile from Marche–Umbria but is not as clearly displayed in this area (Stoll & Schrag 2000).

The lack of exact correlation between the $\delta^{13}\text{C}$ curves illustrated here and those from elsewhere may relate to several factors: stratigraphical gaps in some of the sequences, diagenesis of carbonate in peri-platform ooze and/or organic-rich sediments, whether near-surface or bottom-water chemistry is recorded (planktonic vs benthonic foraminifera) and whether the signal derives from organic carbon or carbonate. With the exception of the negative excursion corresponding to OAE 1b, we are unable to ascertain to what extent the carbon-isotope record from the Coppa della Nuvola section records real changes in the isotopic composition of the dissolved inorganic-carbon reservoir and to what extent these other factors have come into play.

The Coppa della Nuvola $\delta^{18}\text{O}$ curve shows a remarkably similar trend to that of the $\delta^{13}\text{C}$ curve, although the negative excursions are less pronounced. Although clearly influenced by diagenesis, the negative shifts in the lower Albian, upper Albian and lower Cenomanian may be interpreted as a record of warming events. Wilson & Norris (2001) have recently documented, from the upper Albian *R. appenninica* Zone of the western tropical Atlantic Ocean, highly variable sea-water superficial temperatures but with minimum values $4\text{--}6^\circ\text{C}$ higher than the warmest average warm season of today’s tropical oceans. Moreover, the geochemical data from selected planktonic foraminifera suggest a steady thermocline temperature increase during the same stratigraphical interval. The corresponding tract in the Coppa della Nuvola $\delta^{18}\text{O}$ curve, however, records relatively constant values with a very weak positive excursion (Fig. 6). We conclude, given the diagenetic overprint, that secure palaeoceanographic information cannot be unambiguously deciphered from the $\delta^{18}\text{O}$ record of this sequence alone.

6. Calcareous plankton distribution and palaeoenvironmental changes

6a. Planktonic foraminifera

Reference to modern forms, together with stable-isotope data, enables the palaeoecological behaviour, evolution and diversification of Cretaceous planktonic foraminifera to be related to continuing stable oceanic conditions, allowing colonization of the deepest ecological niches by large, complex, keeled forms. By analogy with modern low-latitude oceans, the most diversified Cretaceous planktonic foraminiferal assemblages are expected within a stratified water-column and a deep thermocline (e. g. Hart 1980, 1999; Caron & Homewood 1983; Leckie 1987, 1989; Premoli Silva & Sliter 1999; Leckie et al. 2002). Significant changes in Cretaceous populations are interpreted as reflecting an ecological response to palaeoceanographic variations.

Although the palaeoecological behaviour of Cretaceous planktonic foraminifera is not completely understood, it is possible to place roughly the Albian–Cenomanian genera in an oligotrophic-specialized *versus* eutrophic-opportunist scale. The models generally applied (e.g. Hart 1999; Premoli Silva & Sliter 1999; Leckie et al. 2002), based on palaeoecological evidence and partly on isotopic data, link morphology and size to the depth-habitat. The larger and more complex species were specialized and the deepest-dwelling (e. g., rotaliporids), whereas the smaller and simpler forms (e. g., hedbergellids) inhabited surficial waters and were opportunists (Caron & Homewood 1983; Leckie 1987, 1989; Hart 1999; Premoli Silva & Sliter 1999). Actually, hedbergellids may have thrived in both the thermocline and in surface waters, according to the depth of highest primary productivity, as suggested by the wide range of temperature-related habitats (e. g., Douglas & Savin

[illegible]

with high abundance of radiolarians and may be related to increased eutrophic conditions. Within these intervals, the scarcity of foraminifera could be primary, in part accentuated by the high concentrations of CO₂ related to the eutrophy (caused by oxidation of organic matter in the water column and on the sea floor) which limits carbonate preservation. Pronounced dissolution has been recorded from related strata of the Umbria–Marche basin (Premoli Silva et al. 1989a; Tornaghi et al. 1989, Erba 1992). In the Coppa della Nuvoletta section, these intervals correlate with the main excursions of the carbon- and oxygen-isotope curves (C11–C12, C16–C17) and with the ‘Urbino’ and ‘Amadeus’ events. Both these events, as far as the record of the Gargano area is concerned, appear to be related to episodes of increased productivity.

the Albian time interval. This change of the trophic status of surface water probably caused the decrease in calcareous-nannofossil diversity recorded in the Late Albian–Early Cenomanian interval. Moreover, the two intervals of scarce abundance of planktonic foraminifera, associated with high abundance of radiolarians and with the principal negative $\delta^{13}\text{C}$ excursions, show a decrease in nannofossil diversity which, however, never drops to zero. Within these two intervals, the nannofossil assemblages record a slight increase of the fertility indices *Biscutum constans* and *Zygodiscus* spp. and dominance of *Watznaueria barnesae*. The high abundance of the last species, known as a diagenetic index, is probably related to the previously discussed high concentrations of CO_2 , which can produce pronounced dissolution of carbonate. But, as argued by Negri et al. (2003), *Watznaueria barnesae* was probably also a tolerant, generalized taxon, increasing in abundance during periods of instability in the water masses and high availability of trophic resources in the photic zone.

7. Summary and conclusions

The high-resolution integrated micropaleontological and geochemical analyses of the upper Aptian–lower Cenomanian Coppa della Nuvola section (Gargano Promontory) have provided an opportunity to investigate the record of mid-Cretaceous palaeoceanographic changes. The carbonate carbon- and oxygen-isotope curves represent one of the most detailed records for low latitudes (Fig. 6). The main results of this study are summarized as follows.

1. The Coppa della Nuvola curve displays four negative shifts of $\delta^{13}\text{C}$, respectively in the lower Albian (*T. primula*; *P. columnata* Zones), lower upper Albian (basal *T. praeticinensis* Subzone; *R. achlyostaurion* Zone; weakly developed), upper Albian (upper *T. praeticinensis* Subzone, *R. subticinensis* Zone; *R. achlyostaurion* Zone) followed by marked increases of the values, and in the lower Cenomanian (*R. reicheli* Zone; CC9c Subzone). These excursions are separated by intervals of relatively constant $\delta^{13}\text{C}$ values. The Coppa della Nuvola $\delta^{18}\text{O}$ curve shows a similar trend, albeit with less pronounced excursions.
2. The carbon-isotope curve may be partly correlated with data from Mexico (Bralower et al. 1999) and from the Marche–Umbria Basin (Stoll & Schrag 2000). A reliable correlation exists between the lower Albian successions of the Gargano Promontory, the Vocontian Basin (Southeastern France) and the Mazagan Plateau (Northwest Africa, DSDP Site 545; Herrle 2002) in the record of a marked $\delta^{13}\text{C}$ negative trend, followed by a positive shift, in the *Pre-discosphaera columnata* Zone.
3. The $\delta^{18}\text{O}$ curve records negative shifts in the lower Albian (*T. primula*; *P. columnata* Zones), Upper Albian (upper *T. praeticinensis* Subzone, *R. subticinensis* Zone; *R. achlyostaurion* Zone) and Lower Cenomanian (*R. reicheli* Zone; CC9c Subzone) respectively, which may be interpreted as

records, albeit modified by diagenesis, of warming events. The warming event recorded in the *P. columnata* Zone of the early Albian has been recorded also off northwest Africa and from southern France (Herrle 2002). These $\delta^{18}\text{O}$ signals are interpreted as essentially primary, recording rapid temperature rise consequent upon gas-hydrate release which also resulted in introduction of isotopically light carbon into the global isotopic reservoir.

4. The synchronous negative shifts in $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ may record dissociation of gas hydrates somewhere in the world ocean and the near-instantaneous oxidation of the released methane to another greenhouse gas, carbon dioxide and consequent rapid global warming (Menegatti et al. 1998; Jenkyns 2003). The regional record of the negative isotopic event that precedes or overlaps in time with the regional (Tethys–Atlantic) deposition of black shales characteristic of OAE1b (early Albian Paquier Event) renders this the best-documented example of this phenomenon in the late Aptian–early Cenomanian interval. However, the similar isotopic signatures of the intervals in the upper Albian and lower Cenomanian may record phenomena that were similar in kind.
5. The palaeoceanographic conditions inferred by the biotic and isotopic changes are fluctuating meso-eutrophy through the late Aptian–early middle Albian of the Gargano area. These conditions were related to a weak thermocline, sometimes disrupted by more vigorous vertical mixing. The consequent nutrient recycling caused an increase in surface-water fertility, as suggested by the radiolarian abundance peaks and drop in foraminiferal diversity (Fig. 6). Two main episodes of increased eutrophy occurred in the early Albian and in the early late Albian that correlate with the deposition of black shales in the same stratigraphic interval in nearby areas of the Gargano Promontory and elsewhere. These black shales, although absent in the Coppa della Nuvola section, record the effects of OAE 1b and OAE 1c. More stable oligotrophic conditions, beginning in the late Albian, are documented by biotic signals that record a general marked decrease in abundance of radiolarians and the development of the specialized foraminifera (rotaliporids). However, the planktonic community is characterized by rare and small-sized rotaliporids, suggesting that highly oligotrophic conditions were never attained. The described change in nutrient availability during the Albian probably caused the recorded decrease of calcareous nannofossil diversity.

Acknowledgements

A detailed data set on planktonic foraminifera and calcareous nannofossils (range charts of the studied interval) is available on request. This research was financially supported by MIUR ex 60% to V. L. and FAR to M.C. We acknowledge helpful reviews from Helmut Weissert, Christoph Hemleben and Jürgen Remane.

REFERENCES

- ARTHUR, M. A., DEAN W. E. & PRATT, L. M. 1988: Geochemical and climatic effects of increased marine organic carbon burial at the Cenomanian/Turonian boundary. *Nature* 335, 714–717.
- JENKYN, H. C., BRUMSACK, H. J. & SCHLANGER, S. O. 1990: Stratigraphy, geochemistry, and paleoceanography of organic carbon-rich Cretaceous sequences. In R. N. GINSBURG & B. BEAUDOIN (Eds.): *Cretaceous Resources, Events and Rhythms*, Dordrecht: Kluwer Academic Publishers. 75–119.
- BACCILLE, L., BOSELLINI, A., 1965: Diagrammi per la stima visiva della composizione percentuale nelle rocce sedimentarie: *Annali dell'Università di Ferrara (Nuova Serie), Sezione IX, Scienze Geologiche e Paleontologiche* 1, 117–153.
- BELLANCA, A., CLAPS, M., ERBA, E., MASETTI, D., NERI, R., PREMOLI SILVA, I. & VENEZIA, F. 1996: Orbitally induced limestone/marlstone rhythms in the Albian-Cenomanian Cison section (Venetian region, northern Italy): sedimentology, calcareous and siliceous plankton distribution, elemental and isotope geochemistry. *Palaeogeogr., Palaeoclim., Palaeoecol.* 126, 227–260.
- ERBA, E., NERI, R., PREMOLI SILVA, I., SPROVIERI, M., TREMOLADA, F. & VERGA, D. 2002: Palaeoceanographic significance of the Tethyan 'Livello Selli' (Early Aptian) from the Hybla Formation, northwestern Sicily: biostratigraphy and high-resolution chemostratigraphic records. *Palaeogeogr., Palaeoclim., Palaeoecol.* 185, 175–196.
- BERNOULLI, D. 1972: North Atlantic and Mediterranean Mesozoic facies: a comparison. In: C. D. HOLLISTER & J. I. EWING Eds, *Init. Rep. Deep Sea Drill. Proj.* 11, 801–879.
- BISCHOFF, G. & MUTTERLOSE, J. 1998: Calcareous nannofossils of the Barremian/Aptian boundary interval in NW Europe: biostratigraphic and palaeoecologic implications of a high resolution study. *Cret. Res.* 19, 635–661.
- BOSELLINI, A., NERI, C. & LUCIANI, V. 1993: Platform margin collapse and sequence stratigraphic organization of carbonate slopes: Cretaceous-Eocene, Gargano Promontory, southern Italy. *Terra Nova* 5, 282–297.
- MORSILLI, M. & NERI, C. 1999: Long-term event stratigraphy of the Apulia Platform margin: the Gargano transect (Upper Jurassic to Eocene). *J. Sed. Res.* 69, 1241–1253.
- MORSILLI, M. & NERI, C. 2000: The eastern margin of the Apulia Platform: The Gargano Transect. CRER Working Group 4. Guide Book. Vieste, Gargano, Italy.
- BOWN, P. R., RUTLEDGE, D. C., CRUX, J. A. & GALLAGHER, L. T. 1998: Lower Cretaceous. In: P. R. BOWN (Ed.): *Calcareous Nannofossil Biostratigraphy*, Cambridge, 86–131.
- BRALOWER, T. J. 1987: Valanginian to Aptian calcareous nannofossil stratigraphy and correlation with the upper M-sequence magnetic anomalies. *Mar. Micropal.* 11, 293–310.
- LECKIE, R. M., SLITER, W. V. & THIERSTEIN, H. R. 1995: An integrated Cretaceous microfossil biostratigraphy. In: W. A. BERGGREN, D. V. KENT, M. P. AUBRY & J. HARDENBOL (Eds.). *Geochronology time scales and global stratigraphic correlation*, SEPM Spec. Publ. 54, 56–79.
- COBABBÉ, E., CLEMENT, B., SLITER, W. V., OSBURN, C. L. & LONGORIA, J. 1999: The record of global change in mid-Cretaceous (Barremian-Albian) sections from the Sierra Madre, northeastern Mexico. *J. Foram. Res.* 29, 418–437.
- BRÉHÉRET, J.-G., 1994: The Mid-Cretaceous organic-rich sediments from the Vocontian Zone of the French southeast Basin. In A. MASCLÉ Ed, *Hydrocarbon and Petroleum Geology of France*, Special Publication of the European Association of Petroleum Geologists, Springer-Verlag, Berlin, 295–320.
- CARON, M. & DELAMETTE, M. 1986: Niveaux riches en matière organique dans l'Albien vocontien; quelques caractères du paléoenvironnement; essai d'interprétation génétique. *Doc. Bur. Rech. Géol. Min.* 110, 141–191.
- CARON, M. 1985: Cretaceous planktic foraminifers. In H. M. BOLLI, J. B. SAUNDERS, & K. PERCH-NIELSEN (Eds.): *Plankton Stratigraphy*, Cambridge University Press, 17–86.
- & HOMEWOOD, P. 1983: Evolution of early planktic foraminifers. *Mar. Micropal.* 7, 453–462.
- COBIANCHI, M. & PICOTTI, V. 2001: Sedimentary and biological response to sea-level and paleoceanographic changes of a Lower-Middle Jurassic Tethyan Platform margin (Southern Alps, Italy). *Palaeogeogr., Palaeoclim., Palaeoecol.* 169 (3–4), 219–244.
- LUCIANI, V. & BOSELLINI, A. 1997: Early Cretaceous nannofossils and planktonic foraminifera from northern Gargano (Apulia, southern Italy). *Cret. Res.* 18, 249–293.
- LUCIANI, V. & MENEGATTI, A. 1999: The Selli Level of the Gargano Promontory, Apulia, southern Italy: foraminiferal and calcareous nannofossil data. *Cret. Res.* 20, 255–269.
- COCCIONI, R. 2001: The "Pialli Level" from the latest Albian of the Umbria-Marche Apennines (Italy). *Federazione Italiana di Scienze della Terra, GeolItalia 2001, Settembre 2001, Chieti, abstract*.
- COCCIONI, R. & LUPERTO SINNI, E. 1989: Foraminiferal biostratigraphy of the Vico del Gargano section (Lower Cretaceous) from the Gargano, southern Italy. In R. COCCIONI, S. MONECHI & G. PARISI Eds, *Abstract, 1st Meeting of the Working group 2-Pelagic facies (IGCP Project 262 "Tethyan Cretaceous Correlation")*, Centro stampa, Università di Urbino, 25–27.
- & GALEOTTI, S. 1991: Orbitally induced cycles in benthonic foraminiferal assemblage distribution from Aptian-Albian organic-rich Scisti a Fucoidi (Central Italy) In *Abstract, Pelagic and Flysch facies, 3rd Meeting (IGCP Project 262 "Tethyan Cretaceous Correlation")*, Grenoble, 1991, 1–2.
- CORFIELD, R. M., HALL, M. & BRASIER, M. D., 1990: Stable isotope evidence for foraminiferal habitats during the development of the Cenomanian/Turonian oceanic anoxic event. *Geology* 18, 175–178.
- CREMONINI, G., ELM, C. & SELL, R. 1971: Note illustrative della Carta Geologica d'Italia. Foglio 157, "Monte S. Angelo". Servizio Geologico Italiano, Roma 56 p.
- DAVEY, S. D. & JENKYN, H. 1999: Carbon-isotope stratigraphy of shallow-water limestones and implications for the timing of Late Cretaceous sea-level rise and anoxic events (Cenomanian-Turonian of the peri-Adriatic carbonate platform, Croatia). *Eclogae Geol. Helv.* 92, 163–170.
- DICKENS, G. R., O'NEIL, J. R., REA, D. K. & OWEN, R. M. 1995: Dissociation of oceanic methane hydrate as a cause of the carbon isotope excursion at the end of the Paleocene. *Paleoceanogr.* 10, 965–971.
- DOUGLAS, R. G., & SAVIN, S. M., 1978: Oxygen isotopic evidence for the depth stratification of Tertiary and Cretaceous planktic foraminifera. *Marine Micropaleont.* 3, 175–196.
- ERBA, E. 1988: Aptian-Albian calcareous nannofossil biostratigraphy of the Scisti a Fucoidi cored at Piobbico (central Italy). *Riv. It. Paleont. Strat.* 94, 249–284.
- 1992: Calcareous nannofossil distribution in pelagic rhythmic sediments (Aptian-Albian Piobbico core, central Italy). *Riv. It. Paleont. Strat.* 97, 455–484.
- 1996: The Aptian Stage. In P. F. RAWSON, A. V. DHOND, J. M. HANCOCK & W. J. KENNEDY Eds, *Proceedings of 2nd International Symposium on Cretaceous Stage Boundaries*. *Bull. Inst. Roy. Sci. Nat. Belg.* 66, 31–43.
- COCCIONI, R. & PREMOLI SILVA, I. 1989: Gli Scisti a Fucoidi nell'area umbro-marchigiana: le sezioni della S.S. Apicchiense. In: S. CRESTA, S. MONECHI & G. PARISI (Eds.): *Stratigrafia del Mesozoico e Cenozoico nell'area umbro-marchigiana*, p.146–164. *Memorie descrittive della Carta Geologica d'Italia*, Roma: Istituto Poligrafico e Zecca dello Stato.
- CHANNELL, J. E. T., CLAPS, M., JONES, C., LARSON, R., OPDYKE, B., PREMOLI SILVA, I., RIVA, A., SALVINI, G. & TORRICELLI, S. 1999: Integrated stratigraphy of the Cison Apticore (Southern Alps, Italy): A "reference section" for the Barremian-Aptian interval at low latitudes. *J. Foram. Res.* 29, 371–391.
- ERBACHER, J. 1994: Entwicklung und Paläoceanographie mittelkretazischer Radiolarien der westlichen Tethys (Italien) und des Nordatlantiks. *Tübinger Mikropaläont. Mitt.* 12, p. 120.
- ERBACHER, J. & THUROW, J. 1997: Influence of anoxic events on the evolution of mid-Cretaceous radiolaria in the North Atlantic and western Tethys. *Mar. Micropal.* 30, 139–158.
- HEMLEBEN, C., HUBER, B. T. & MARKEY, M. 1999: Correlating environmental changes during Early Albian oceanic anoxic event 1B using benthic foraminiferal paleoecology. *Mar. Micropal.* 38, 7–28.

- FASSELL, M. L., and BRALOWER, T. J. 1999: A warm, equable mid-Cretaceous: Stable isotopic evidence. In: BARRERA, E., & JOHNSON, C., (Eds.): The evolution of Cretaceous Oceanic/Climate Systems, Geological Society of America Special Publication, Geol. Soc. Am. Spec. Pap., 332, 121–142.
- FERRERI, V., WEISSERT, H., D'ARGENIO, B. & BUONCUNTO, F.P. 1997. Carbon isotope stratigraphy: a tool for basin to carbonate platform correlation. *Terra Nova* 9, 57–61.
- GALE, A. S., KENNEDY, W. J., BURNETT, J. A., CARON, M. & KIDD, B. E. 1996: The Late Albian to Early Cenomanian succession at Mont Risou near Rosans (Drome, SE France): an integrated study (ammonites, inoceramids, planktonic foraminifera, nannofossils, oxygen and carbon isotopes). *Cret. Res.* 17, 515–606.
- GRAZIANO, R. 2000: The Aptian-Albian of the Apulia carbonate platform (Gargano Promontory, southern Italy): evidence of palaeoceanographic and tectonic controls on the stratigraphic architecture of the platform margin. *Cret. Res.* 21, 107–126.
- GRÖTSCH, J., BILLING, I. & VAHRENKAMP, V. 1998: Carbon-isotope stratigraphy in shallow-water carbonates: implications for Cretaceous black-shale deposition. *Sedimentology* 45, 623–634.
- HALLOCK, P. 1987: Fluctuations in the trophic resource continuum: a factor in global diversity cycles? *Paleoceanography* 2, 457–471.
- HARDENBOL, J., THIERRY, J., FARLEY M.B., JACQUIN, T., DE GRACIANSKY, P.C. & VAIL, P. 1998: Mesozoic and Cenozoic Sequence Chronostratigraphic framework of European Basins In P.C. DE GRACIANSKY, J. HARDENBOL, T. JACQUIN, & P.R. VAIL, Eds. *Mesozoic and Cenozoic sequence stratigraphy of European basins*. SEPM Spec. Publ. 60, 763–781.
- HART, M.B. 1980: A water depth model for the evolution of the planktonic foraminifera. *Nature* 268, 252–254.
- 1999: The evolution and biodiversity of Cretaceous planktonic foraminifera. *Geobios* 32, 247–255.
- HESELBO, S. P., GRÖCKE, D. R., JENKINS, H. C., BJERRUM, C. J., FARRIMOND, P., MORGANS BELL, H. S. & GREEN O. 2000: Massive dissociation of gas hydrate during a Jurassic oceanic anoxic event. *Nature* 406, 392–395.
- HERRLE, J. O. 2002: Paleoclimatic and paleoceanographic implications on Mid-Cretaceous black-shale formation in the Vocontian Basin and the Atlantic: evidence from calcareous nannofossils and stable isotopes. *Tübinger Micropalaontologische Mitteilungen* 27, p. 114.
- HILL M. E. 1976: Lower Cretaceous calcareous nannofossils from Texas and Oklahoma. *Palaeontographica B.*, 156, 103–179.
- JAHREN, A. H., ARENS, N. C., SARMENTO, G., GUERRIERO, J. & AMUNDSON, R. 2001: Terrestrial record of methane hydrate dissociation in Early Cretaceous. *Geology* 29, 159–162.
- JENKINS, H. C. 1980: Cretaceous anoxic events: from continents to oceans. *Journ. Geol. Soc. London* 137, 137–188.
- 1995: Carbon-isotope stratigraphy and paleoceanographic significance of the Lower Cretaceous shallow-water carbonates of Resolution Guyot, Mid-Pacific Mountains. *Proc. Ocean. Drill. Prog., Scientific Results* 143, 99–105.
- 1999: Mesozoic anoxic events and palaeoclimate. *Zentr. Geol. Palaont.* 1997, 943–949.
- 2003: Evidence for rapid climate change in the Mesozoic–Palaeogene greenhouse world. *Phil. Trans Roy. Soc., Ser. A* 361, 1885–1916.
- & WILSON, P. A. 1999: Stratigraphy, paleoceanography, and evolution of Cretaceous Pacific guyots: relics from a greenhouse Earth. *Am. Jour. Sci.* 299, 341–392.
- GALE, A.S. & CORFIELD, R. M. 1994: Carbon- and oxygen-isotope stratigraphy of the English Chalk and Italian Scaglia and its palaeoclimatic significance. *Geol. Mag.* 131, 1–34.
- JONES, C.E. & JENKINS, H.C. 2001: Seawater strontium isotopes, oceanic anoxic events, and seafloor hydrothermal activity in the Jurassic and Cretaceous. *Am. Jour. Sci.* 301, 112–149.
- KENNEDY, W.J., GALE, A.S., BOWN, P.R., CARON, M., DAVEY, R.J., GRÖCKE, D. & WRAY, D.S. 2000: Integrated stratigraphy across the Aptian-Albian boundary in the Marnes Bleues, at the Col de Pre-Guittard, Arnyon (Drome), and at Tartonne (Alpes-de-Haute-Provence), France: a candidate Global Boundary Stratotype section and Boundary Point for the base of the Albian Stage. *Cret. Res.* 21, 591–720.
- KUYPERS, M. M. M., BLOKKER, P., ERBACHER, J., KINKEL, H., PANCOST, R.P., SCHOUTEN, S. & DAMSTÉ, J.S.S. 2001: Massive expansion of marine Archea during a mid-Cretaceous oceanic anoxic event. *Science* 293, 92–94.
- LARSON, R. L. 1991: Latest pulse of Earth: Evidence for a mid-Cretaceous superplume. *Geology* 19, 547–550.
- & ERBA, E. 1999: Onset of the mid-Cretaceous greenhouse in the Barremian-Aptian: Igneous events and the biological, sedimentary, and geochemical responses. *Paleoceanography* 14, 663–678.
- LECKIE, R. M. 1987: Paleocology of mid-Cretaceous planktonic foraminifera: a comparison of open ocean and epicontinental sea assemblages. *Micropalaentology* 33, 164–176.
- 1989: A paleoceanographic model for the early evolutionary history of planktonic foraminifera. *Palaeogeogr., Palaeoclim., Palaeoecol.* 73, 107–138.
- BRALOWER, T.J. & CASHMAN, R. 2002: Oceanic anoxic events and plankton evolution: Biotic response to tectonic forcing during the mid-Cretaceous. *Paleoceanography* 17, 10.1029/2001/PA000623.
- LUCIANI, V. 1993: Mid-Turonian planktic foraminifera and depth stratification along a carbonate paleoslope (Gargano Promontory, southern Italy). *An. Univ. Ferrara, Nuova Serie, Sezione: Scienze della Terra* 4, 47–79.
- & COBIANCHI, M. 1994: Type section of the Mattinata Formation (Lower Cretaceous, Gargano Promontory, southern Italy): new biostratigraphic data (calcareous nannofossils and planktonic foraminifers). *Mem. Sci. Geol. Univ. Padova* 46, 283–301.
- & COBIANCHI, M. 1999: The Bonarelli Level and other black shales in the Cenomanian-Turonian of the northeastern Dolomites (Italy): calcareous nannofossil and foraminiferal data. *Cret. Res.* 20, 135–167.
- COBIANCHI, M. & JENKINS, H. C. 2001: Biotic and geochemical response to anoxic events: the Aptian pelagic succession of the Gargano Promontory (southern Italy). *Geol. Mag.* 137, 277–298.
- LUPERTO SINNI, E. 1996: Schema stratigrafico del Cretacico del Gargano basato su risultati di recenti ricerche. *Mem. Soc. Geol. It.* 51, 1019–1036.
- & MASSE, J.P. 1987: Données nouvelles sur la stratigraphie et la micropaléontologie des séries carbonatées de talus et de bassin du Crétacé inférieur du Gargano (Italie Méridionale). *Riv. It. Paleont. Strat.* 93, 347–378.
- & BORGOMANO, J. 1994: Stratigrafia del Cretaceo superiore in facies di scarpata di Monte Sant'Angelo (promontorio del Gargano, Italia Meridionale). *Boll. Soc. Geol. It.* 113, 355–382.
- MANIVIT, H., PERCH-NIELSEN, K., PRINS, B. & VERBEEK, J.W. 1977: Mid-Cretaceous calcareous nannofossil biostratigraphy. *Proc. Kon. Nederl. Akad. Wet. B* 80, 169–181.
- MARSHALL, J.D. 1992: Climatic and oceanographic isotopic signals from the carbonate rock record and their preservation. *Geol. Mag.* 129, 143–160.
- MARTINIS, B. & PAVAN, G. 1967: Note illustrative della Carta Geologica d'Italia. Foglio 157 "Monte S. Angelo", Servizio Geologico Italiano, 56 pp.
- MENEGATTI, A.P., WEISSERT, H., BROWN, R.S., TYSON, R.V., FARRIMOND, P., STRASSER, A. & CARON, M. 1998: High-resolution $\delta^{13}\text{C}$ stratigraphy through the early Aptian "Livello Selli" of the Alpine Tethys. *Paleoceanography* 13, 530–545.
- MERLA, G., ERCOLI, A. & TORRE, D. 1969: Note illustrative della carta geologica d'Italia alla scala di 1:100.000, Foglio 164 "Foggia". Servizio Geologico d'Italia, 22 pp.
- MITCHELL, S. F., PAUL, C. R. C. & GALE, A. S. 1996: Carbon isotopes and sequence stratigraphy. In: J. A. HOWELL & J. F. AITKEN (Eds.): *High Resolution Sequence Stratigraphy: Innovations and Applications*. Spec. Publ. Geol. Soc. London 104, 11–24.
- MORSILLI, M. 1998: Stratigrafia e sedimentologia del margine della Piattaforma Apula nel Gargano (Giurassico superiore-Cretacico inferiore). PhD thesis, Università di Ferrara, National Libraries of Rome and Florence, 203 pp.
- BOSELLINI, 1997: Carbonate facies zonation of the Upper Jurassic-Lower Cretaceous Apulia Platform margin (Gargano Promontory, southern Italy). *Riv. It. Paleontol. Strat.* 103, 193–206.
- MUTTERLOSE, J. 1998: The Barremian-Aptian turnover of biota in north-western Europe: evidence from belemnites. *Palaeogeogr., Palaeoclim., Palaeoecol.* 144, 161–173.

- NEGRI, A., COBIANCHI, M., LUCIANI, V., FRABONI, R., MILANI, A. & CLAPS, M. 2003: Tethyan Cenomanian pelagic rhythmic sedimentation and Pleistocene sapropels: is the biotic signal comparable? *Palaeogeography, Palaeoclimatology, Palaeoecology*, 190, 373–397.
- NERI, C. & LUCIANI, V. 1994: The Monte S. Angelo Sequence (Late Cretaceous–Paleocene, Gargano Promontory, southern Italy): physical stratigraphy and biostratigraphy. *Giornale di Geologia* 52, 149–165.
- NORRIS, R. D., & WILSON, P. A., 1998: Low-latitude sea-surface temperature for the mid-Cretaceous and the evolution of planktonic foraminifera. *Geology* 26, 823–826.
- & WILSON, P. A. 1999: Low-latitude sea-surface temperature for the mid-Cretaceous and the evolution of planktonic foraminifera: Comment and Reply. *Geology* 27, 857.
- PADDEN, M., WEISSERT, H. & DE RAFELIS, M. 2001: Evidence for Late Jurassic methane hydrate release: opening of a Pangean gateway. *Geology*, 29, 223–226.
- PATTERSON, W. & WALTER, L. M. 1994: Depletion of ^{13}C in seawater ΣCO_2 on modern carbonate platforms: significance for the carbon isotopic record of carbonates. *Geology* 22, 885–888.
- PAVAN, G. & PIRINI, C. 1966: *Stratigrafia del Foglio 157 "Monte S. Angelo"*. Boll. Serv. Geol. It. 86, 123–189.
- PERCH-NIELSEN, K. 1985: Mesozoic calcareous nannofossils In: H.M. BOLL, J.B. SAUNDERS & K. PERCH-NIELSEN Eds, *Plankton Stratigraphy*. Cambridge University Press, 329–426.
- PREMOLI SILVA, I. & SLITER, W.V. 1994: Cretaceous planktonic foraminiferal biostratigraphy and evolutionary trends from the Bottaccione section, Gubbio (Italy). *Palaeontographia Italica* 82, 1–89.
- & SLITER, W.V. 1999: Cretaceous paleoceanography: evidence from planktonic foraminiferal evolution. *Geol. Soc. Am. Spec. Paper* 332, 301–328.
- ERBA, E. & TORNAGHI, M.E. 1989a: Paleoenvironmental signals and changes in surface fertility in mid Cretaceous Corg-rich facies of the Fucoid Marls (central Italy). *Géobios Mém. Spéc.* 11, 225–236.
- RIPEPE, M. & TORNAGHI, M. E. 1989b: Planktonic foraminiferal distribution record productivity cycles: evidence from the Aptian-Albian Piobbico core (central Italy). *Terra Nova* 1, 443–448.
- ERBA, E., SALVINI, G., LOCATELLI, C. & VERGA, D. 1999: Biotic changes in Cretaceous oceanic anoxic events of the Tethys. *J. Foram. Res.* 29, 352–370.
- PRICE, G. D., and HART, M. B., 2002: Isotopic evidence for Early to mid-Cretaceous ocean temperature variability: Marine Micropaleontology 46, 45–58.
- ROBASZYNKI, F. & CARON, M. 1995: Foraminifères planctoniques du Crétacé: commentaire de la zonation Europe-Méditerranée. *Bull. Soc. Géol. France* 166, 691–692.
- RÖHL, U., BRALOWER, T. J., NORRIS, R. D. & WEFER, G. 2000. New chronology for the late Paleocene thermal maximum and its environmental implications. *Geology*, 28, 927–930.
- ROTH, P. H. 1983: Jurassic and Lower Cretaceous calcareous nannofossils in the western Atlantic (Site 534). *Biostratigraphy, preservation and some observations on biogeography and paleoceanography*. *Init. Rep. Deep Sea Drill. Proj.* 76, 587–621.
- 1987: Mesozoic calcareous nannofossil evolution: relation to paleoceanographic events. *Paleoceanography* 2, 601–611.
- 1989: Ocean circulation and calcareous nannoplankton evolution during the Jurassic and Cretaceous. *Palaeogeogr. Palaeoclim. Palaeoecol.* 74, 111–126.
- SCHLAGER, W. & JAMES, N. P. 1978: Low-magnesian calcite limestones forming at the deep-sea floor, Tongue of the Ocean, Bahamas. *Sedimentology*, 25, 675–702.
- SCHLANGER, S. O. & JENKINS, H. C. 1976: Cretaceous oceanic anoxic events: causes and consequences. *Geo. Mij.* 55, 179–184.
- ARTHUR, M. A., JENKINS, H. C. & SCHOLLE, P. A. 1987: The Cenomanian-Turonian oceanic anoxic event. I. Stratigraphy and distribution of organic carbon-rich beds and the marine $\delta^{13}\text{C}$ excursion. In: J. BROOKS, J. & A. J. FLEET, (Eds.): *Marine Petroleum Source Rocks*. *Geol. Soc. London, Spec. Publ.* 26, 371–399.
- SCHOLLE, P. & ARTHUR, M. A. 1980: Carbon isotopic fluctuations in pelagic limestones: Potential stratigraphic and petroleum exploration tool. *AAPG Bull.* 64, 67–87.
- SIGAL, J. 1977: Essai de zonation du Crétacé méditerranéen à l'aide des foraminifères planctoniques. *Géol. Méditer.* 4, 99–108.
- SISSINGH, W. 1977: Biostratigraphy of Cretaceous calcareous nannoplankton. *Geol. Mij.* 56, 37–65.
- SLITER, W.V. 1989a: Biostratigraphic zonation for Cretaceous planktonic foraminifera examined in thin section. *J. Foram. Res.* 19, 1–19.
- 1989b: Aptian anoxia in the Pacific Basin. *Geology* 17, 909–912.
- SLITER, W. V. 1992: Cretaceous planktonic foraminiferal biostratigraphy and paleoceanographic events in the Pacific Ocean with emphasis on indurated sediment. In: K. ISHIZAKI & T. SAITO, (Eds.): *Centenary of Japanese Micropaleontology*. Terra Scientific Publishing Co., 281–299.
- STOLL, H. M. & SCHRAG, D. P. 2000: High resolution stable isotope records from the Upper Cretaceous rocks of Italy and Spain: Glacial episode in a greenhouse planet? *Geol. Soc. Am. Bull.* 112 (2), 308–319.
- STRASSER, A., CARON, M. & GJERMENI, M. 2001: The Aptian, Albian and Cenomanian of Roter Sattel, Romandes Prealps, Switzerland: a high-resolution record of oceanographic changes. *Cret. Res.* 22, 183–199.
- TARDUNO, J. A., SLITER, W. V., KROENKE, L., LECKIE, M., MAYER, H., MAHONEY, J. J., MUSGRAVE, R., STOREY, M. & WINTERER, E. L. 1991: Rapid formation of the Ontong Java Plateau by Aptian mantle plume volcanism. *Science* 18, 399–403.
- THIERSTEIN, H. R. 1971: Tentative Lower Cretaceous calcareous nannoplankton zonation. *Eclogae Geol. Helv.* 64, 459–488.
- THIERSTEIN, H. R. 1973: Lower Cretaceous calcareous nannoplankton biostratigraphy. *Abh. Geol. Bundesanst.* 29, 1–52.
- THIERSTEIN, H. R. 1976: Mesozoic calcareous nannoplankton biostratigraphy of marine sediments. *Mar. Micropal.* 1, 325–362.
- TORNAGHI, M. E., PREMOLI SILVA, I., & RIPEPE, M. 1989: Lithostratigraphy and planktonic foraminiferal biostratigraphy of the Aptian-Albian "Scisti a Fucoidi" in the Piobbico core, Marche, Italy: background for cyclostratigraphy. *Riv. It. Paleont. Strat.* 95 (3), 223–264.
- VAN HINTE, J.E. 1976: A Cretaceous time scale. *AAPG Bull.* 60, 498–516.
- WILSON, P.A. & NORRIS, R.D. 2001: Warm tropical ocean surface and global anoxia during the mid-Cretaceous period. *Nature* 412, 425–429.
- ZAPPATERA, E. 1990: Carbonate paleogeographic sequence of the periadriatic region. *Boll. Soc. Geol. It.* 109, 5–20.

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