Zeitschrift: Ingénieurs et architectes suisses

Band: 116 (1990)

Heft: 18

Artikel: Impact of Weoclawek reservoir on hydromorphological changes of the

lower vistula channel downstream from the dam (Poland)

Autor: Babinski, Z.

DOI: https://doi.org/10.5169/seals-77297

Nutzungsbedingungen

Die ETH-Bibliothek ist die Anbieterin der digitalisierten Zeitschriften auf E-Periodica. Sie besitzt keine Urheberrechte an den Zeitschriften und ist nicht verantwortlich für deren Inhalte. Die Rechte liegen in der Regel bei den Herausgebern beziehungsweise den externen Rechteinhabern. Das Veröffentlichen von Bildern in Print- und Online-Publikationen sowie auf Social Media-Kanälen oder Webseiten ist nur mit vorheriger Genehmigung der Rechteinhaber erlaubt. Mehr erfahren

Conditions d'utilisation

L'ETH Library est le fournisseur des revues numérisées. Elle ne détient aucun droit d'auteur sur les revues et n'est pas responsable de leur contenu. En règle générale, les droits sont détenus par les éditeurs ou les détenteurs de droits externes. La reproduction d'images dans des publications imprimées ou en ligne ainsi que sur des canaux de médias sociaux ou des sites web n'est autorisée qu'avec l'accord préalable des détenteurs des droits. En savoir plus

Terms of use

The ETH Library is the provider of the digitised journals. It does not own any copyrights to the journals and is not responsible for their content. The rights usually lie with the publishers or the external rights holders. Publishing images in print and online publications, as well as on social media channels or websites, is only permitted with the prior consent of the rights holders. Find out more

Download PDF: 06.08.2025

ETH-Bibliothek Zürich, E-Periodica, https://www.e-periodica.ch

IMPACT OF WEOCLAWEK RESERVOIR ON HYOROMORPHOLOGICAL CHANGES OF THE LOWER VISTULA CHANNEL DOWNSTREAM FROM THE DAM (POLAND)

Z. BABINSKI

Institute of Geography and Spatial Organization Polish Academy of Sciences, Department of Geomorphology and Hydrology of Lowland. 87-100 Torun, Kopernika 19, Poland.

ABSTRACT The contribution deals with problems of channel processes of the Vistula River belLow the Wloclawek dam. The rate of development of the erosional zone and the volume of the material eroded from the channel bed during 19 years after the construction of the dam has been discussed. Moreover, the consequences of bed erosion for the environmental changes near to the erosional zone and downstream of it were presented.

INTRODUCTION

The effects of Water dams on channel processes in lower reaches have been presented by Veksler and Donenberg (1983) on examples of 10 reservoirs in USRR and by Williams and Wolman (1984) on examples of 21 water reservoirs in USA. The results obtained by these authors as well as data collected by Raynov et al. (1986) indicate that the development of the erosional zone is multi-stages and highly differentiated in space. In general, it is believed that bed erosion is most intense at the beginning of the reservoir's activity, when it extent might reach 120 km downstream from the dam and vertical incision -7.5 m, locally even 19.0 m (Raynov et al. 1986). Wlocławek reservoir represented average extent of water storage.

CHARACTERISTIC OF WLOCLAWEK RESERVOIR

The Wloclawek dam is the first and up to now the sole structure of that type in the designed lower Vistula cascade. The dam was located on the 675 km of the Vistula River Course. Ilt closes river basin of an area of about 171 000 km2, which is equal to 55% of the whole territory of Poland. The dam construction resulted in water elevation by 10.7 m above the mean annual water stage of the river. At the normal level, the water reservoir occupied an area of about 70 km² and it has 408 millions m³ of the total storage capacity. The mean discharge values of Vistula is about 980 m3s. The biologically inviolable discharge is 350 m^3s^{-1} and $Q = 0.3\% - 10600 \text{ m}^3\text{s}^{-1}$.

BED EROSION DOWNSTREAM FROM THE DAM

The partition of the Vistula channel by the Wloclawek dam which occurred in October 1968, led to changes in the development of fluvial processes t especially in bed erosion (Babiriski, 1982). The reasons of this process are: stoppage of bed load transport which is accumulated in the upper portion of the reservoir and in temporary increases of the river energy connected with twenty-four hours oscillations of water level which reached maximum three metres. These fluctuations of water level are still percep-

tible at the distance of 200 km downstream from the dam (Babiriski 1982).

The process of the bed erosion downstream from the dam was not uniform but variable in time and space. It was moving downstream as an "erosional wave", characterized by the balance of material eroded from bottom ($R_{\rm E}$) and by the rate of the front of the erosional zone shifting (Figs 1, 2).

Baseing on these data, constructed the straight regression until 2000, for a prognosin of the channel bed degradation downstream from the dam till the ending of the new flood plain creating (Fig. 2).

Since a four year period after the damming, erosion effects were observed at a distance of 9 km-downstream from the dam, whereas fifteen years later on, erosion exerted its influence as far as nearly 24 km (Fig. 2). Over 4 millions m³ of the bed material were eroded in the former case, whereas 14.6 million m³ of sand were removed in the later, case. In the first 4 years, the average rate of the bed erosion was 1 million m³ year-¹ by the shifting of the front of the erosional zone 2.25 km year-¹, in the following 12 years volume of the removed material dropped to about 0.5 million m³ year-¹, by

an average rate of shifting of the erosional zone 0.58 km year ¹ (Fig. 1). Thus, there is a general decrease of the erosional activity of the flowing water.

In a quite unexpected way proceeded the processes of bed erosion during the last 3 years, 1984-1987 (Fig. 1). In that time, despite of the decreasing water discharge ("dry" years), the shifting rate of the front of the erosional zone increased up to 2.7 km year-1 by the rate of erosion 1.6 million m3 year-1, i.e. a substantial value greater than at the beginning of the activity of the reservoir (Fig. 1). The differentiation of the rate of the channel processes should be explained by this circumstance that in the first 4 years an intense exchange of the river load proceeded on the way: channel floor flood plain being formed. Afterwards, in spite of similar conditions, the rate of erosion decreased which was associated with the lengthening of the erosional zone and the incision of the channel to the resistant boulder pavement and Tertiary clays. In addition, in the last period, the exclusion of the newly formed flood plain from the channel processes, resulted in the limitation of the erosion only to the watercurrent zone (thalweg) about 0.4 km wide,

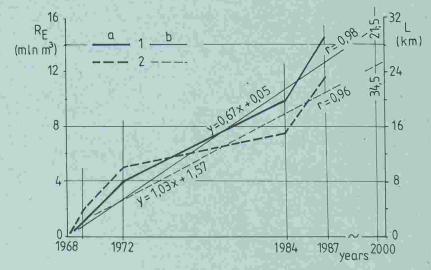


Fig. 1. The balance of the bed erosion $-R_{\rm E}$ (la) and the rate of the front of erosional zone shifting -L (2a) downstream from the Wloclawek dam between 1968-1987, and it's straight regression until 2000 (1b and 2b respectively).

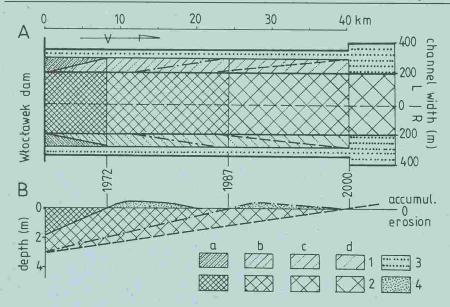


Fig. 2. The rate of channel changes downstream from the Wloclawek dam in a plane (A) and cross section (B). 1 – the stages of the new flood plain development, 2 – the stages of erosional zone development: a/until 1972, b/until 1987, c/until 2000, d/until 2020, when the process of new flood plain creation will be ending, 3 – the flood plain created by regulation works, 4 – accumulation zone – braided channel.

stretching up to 23.7 km downstream from the dam (Fig. 2). This process, determined by the deepening of the erosional zone and the creation of a new flood plain is going to a new type of channel formation from braided to straight one.

ACCUMULATION ZONES CREATED BY EROSIONAL PROCESS

The sand and gravel material eroded from the channel bottom downstream from the dam was deposited in its main part below the erosional zone, creating typical braided channel with side- and central bars (Fig. 2). Moreover, a part of the eroded material was deposited in the bank zone between groynes. This fact together with the constant tendency to lowering of the channel bottom resulted in rising and fixing lateral bars. In such way a new flood plain was formed with the surface by 0.5-2.0 m lower than that of the former one, and in the sector of 8 km long after the 4 years of the reservoir existing and more than 20 km long after nineteen years (Fig. 2). This

process of the new flood plain formation should be ended (as was computed from straight regression - Figure 1), until 2020.

CONSEQUENCES OF BED EROSION DOWNSTREAM FROM THE DAM

As a result of the deepening of the channel bottom below the Wloclawek dam were changes in hydromorphological conditions of Vistula River. The effects of this process are lessening of channel bed slope and lowering of water table level along that reach. This fact resulted also in changes of the grain-size composition of the bed load material. The new hydrodynamical conditions influenced on the pattern of the channel and also contributed to changes in the system of channel mesoforms. Typical braided river contained With numerous islands, central- and linguoid bars changed to straight and limited of groynes meandering channel with diagonal bars together with pools occurring in alternation prevail.

REFERENCES

Babiriski, Z. (1982) Procesy korytowe Wisly ponizej zapory wodnej we Wlocławku ("Fluvial processes of the Vistula River downstream from the Wlocławek dam). Dokumentacja Geograf. 1-2, IG PAN

Raynov, S., Pechinov, D., & Kopaljani, Z. (1986) River response to hydraulic structures. International Hydrological Programme, UNESCO, Paris.

Veksler, A.B. & Donenberg, V.M. (1983) Pereformirovanija rusla v nizhnikh befakh krupnykh gidroelektrostancyj (Channel changes below great power stations), Energoatomizdat.

Williams, G.P. & Wolman, M.G. (1984) Downstream effects of dams on alluvial rivers, Geolog. Surv. Prof. Paper, 1286, Washington

CONSÉQUENCES DE LA MODIFICATION DU RÉGIME D'ÉCOULEMENT A L'AVAL DES RETENUES

D. DUBAND

Electricité de France – Division Technique Générale 37, Rue Diderot BP 41 – Centre de tri 38040 Grenoble Cédex – France

RESUME Les retenues ont une incidence reconnue, prévue et surveillée sur le régime d'écoulement des rivières à l'aval. L'évolution des usages de l'eau par les hommes dans un cadre climatique non stationnaire et le souci d'un équilibre écologique font que ces retenues se révèlent précieuses et que l'on envisage d'élargir le champ de leur gestion, qui concernait essentiellement jusqu'à présent, la production de kilowatt heures au moindre coût en France. Les éléments et esquisses existent pour élaborer une économie de l'eau adaptée et cohérente.

INTRODUCTION

Construire des retenues sur des rivières ou à proximité, c'est se donner les moyens de stocker l'eau pendant un certain temps pour la redistribuer. On opère un transfert en volume d'eau d'une période de l'année pendant laquelle il y a abondance à une autre où il y a pénurie, ou d'une année à

l'autre pour satisfaire des besoins et usages nationaux - régionaux - locaux qui sont parfois antagonistes dans la gestion des ressources en eau, en lissant ainsi les aléas saisonniers et climatiques de la météorologie (pluie, température...):

- production d'énergie électrique (con-

sommation industrielle et domestique)

- irrigation (agriculture)

soutien des étiages (agriculture - eau potable)

- alimentation en eau potable

- écrêtement des crues (sécurité)

- loisirs nautiques (tourisme)